

## Reference:

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## Heat and Water transfer in Soils

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The processes of heat and water transfer in soils influence the boundary conditions for heat and water flow at the lower limit of the atmosphere. These processes are therefore of importance in the study of both the microclimate and macroclimate (energy balance, transformation of air masses). The experimental and theoretical investigation of these phenomena has advanced rapidly during the past ten years.

Water moves in the soil in the liquid and in the vapour phase. Liquid transfer is governed by Darcy's law (Baver, 1956), which can be written:

$$q_l = -\rho_l K \nabla \Phi = -\rho_l K \nabla \psi - \rho_l K \mathbf{k} \quad (1)$$

where  $q_l$  ( $\text{g cm}^{-2} \text{sec}^{-1}$ ) is the liquid flux density,  $\rho_l$  ( $\text{g cm}^{-3}$ ) the density of liquid water,  $\Phi$  (cm) the hydraulic head and  $K$  ( $\text{cm sec}^{-1}$ ) the hydraulic conductivity.  $\Phi = \psi + z$  where  $\psi$  (cm) is the pressure head and  $z$  (cm) the gravitational head. If atmospheric pressure is taken as the datum for  $\psi$  this quantity is positive below the water table and negative above the water table.  $\mathbf{k}$  is a unit vector in the  $z$ -direction (positive upwards). In unsaturated soil both  $\psi$  and  $K$  decrease rapidly with decreasing moisture content.

Movement in the vapour phase is a process of diffusion of water vapour in the air-filled pores. The vapour flux density,  $q_v$  ( $\text{g cm}^{-2} \text{sec}^{-1}$ ) can therefore be expressed as:

$$q_v = -D_v \nabla p_v \quad (2)$$

where  $\rho_v$  ( $\text{g cm}^{-3}$ ) is the vapour density and  $D_v$  ( $\text{cm}^2 \text{sec}^{-1}$ ) a generalized diffusion coefficient (Philip and de Vries, 1957). The vapour density can be written as:

$$\rho_v = \rho_0 h = \rho_0 \exp(g\gamma/RT) \quad (3)$$

where  $\rho_0$  is the saturation vapour density and  $h$  the relative humidity of the soil air. The latter is related to  $\gamma$  and therefore also to the moisture content.  $g$  ( $\text{cm}^2 \text{sec}^{-1}$ ) is the acceleration of gravity,  $R$  ( $\text{erg g}^{-1} \text{OK}^{-1}$ ) the gas constant for 1 g of water vapour and  $T$  ( $^{\circ}\text{K}$ ) the absolute temperature.

Heat transfer in soils is primarily a process of heat conduction, but water movement also contributes to the heat flow mainly by transfer of latent heat (de Vries, 1952). The heat flux density  $q_h$  ( $\text{cal cm}^{-2} \text{sec}^{-1}$ ) can be expressed as follows:

$$q_h = -\lambda \nabla T + L q_v \quad (4)$$

where  $\lambda$  ( $\text{cal cm}^{-1} \text{sec}^{-1} \text{OK}^{-1}$ ) is the thermal conductivity and  $L$  ( $\text{cal g}^{-1}$ ) the heat of vaporization of water.

From equations (1) to (4) simultaneous differential equations can be derived with the moisture content and the temperature as independent variables (Philip and de Vries, 1957). These equations are non-linear and for their solution one must have recourse to numerical methods.

Infiltration is mainly governed by moisture movement in the liquid phase under the influence of gravity and gradients of moisture content (Philip 1957 a,b). So is evaporation from bare soil at evaporation rates on the order of  $10^{-6} \text{ cm sec}^{-1}$  or more (Philip, 1957 c; de Vries 1957). Under natural conditions vapour flow in the soil due to temperature gradients is of the order of  $10^{-7} \text{ cm sec}^{-1}$  or less.

For further meteorological and hydrological implications of these principles the reader is referred to the literature cited below, which also contains many references to earlier work.

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